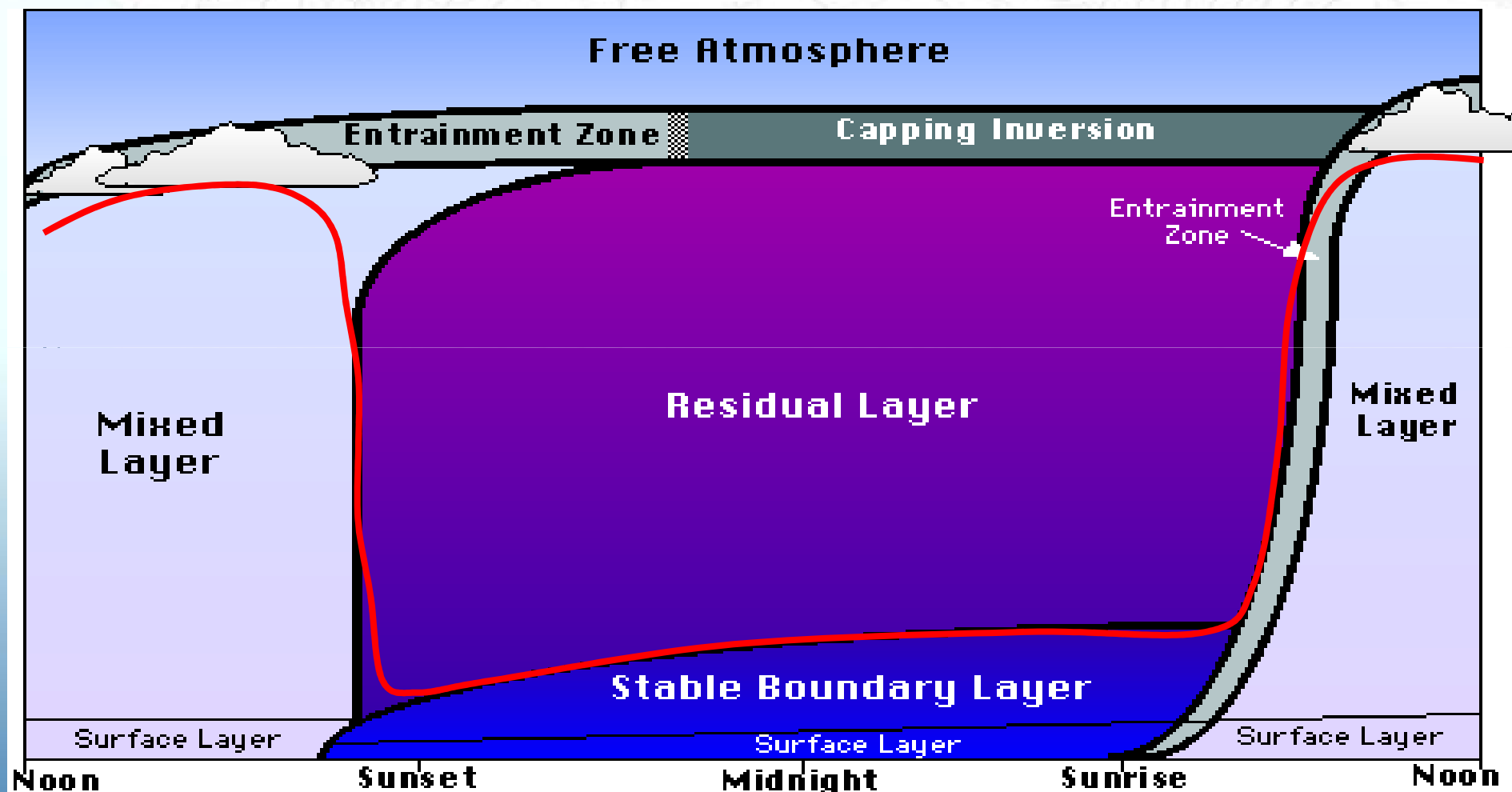


# MLH from MWR

## a quick review

Nico Cimini, CNR-IMAA

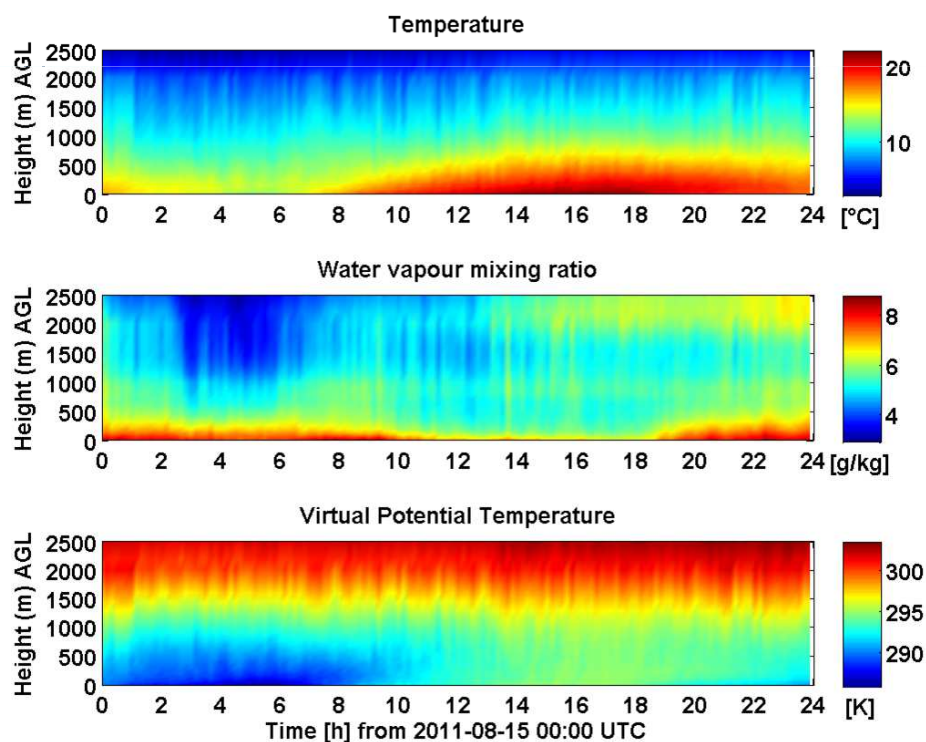


## Structure of Planetary Boundary Layer

*adapted from Stull, 1988.*

# What ground-based MWR provide?

- Ground-based MWR provide T and H profiles
  - High temporal resolution (~1 min),
  - Low-to-moderate vertical resolution,
  - Information mostly residing in the PBL



# Mixing Layer Height (MLH) retrievals

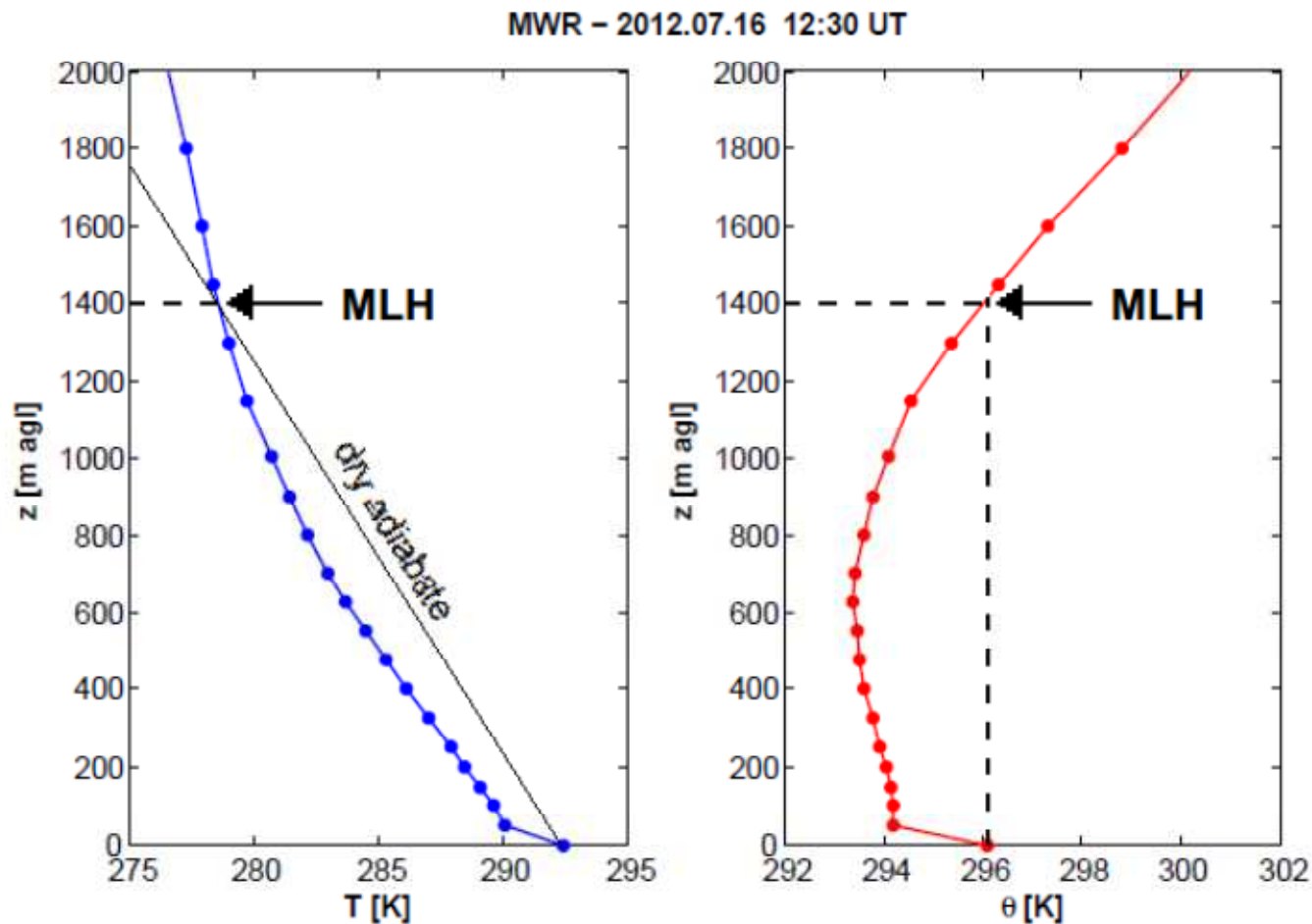
## Methods based on $T$ and $H$ profiles\*

- Parcel Method (PM)
- Surface-Based Temperature Inversion (SBTI)
- Gradient Method (*Liu and Liang Method*)

\*Collaud coen et al., 2014; Seibert et al., 2000; Seidel et al., 2010.

# Parcel Method (PM)

PM\* defines the MLH as the elevation to which an air parcel with ambient surface  $T$  can rise adiabatically from the ground.



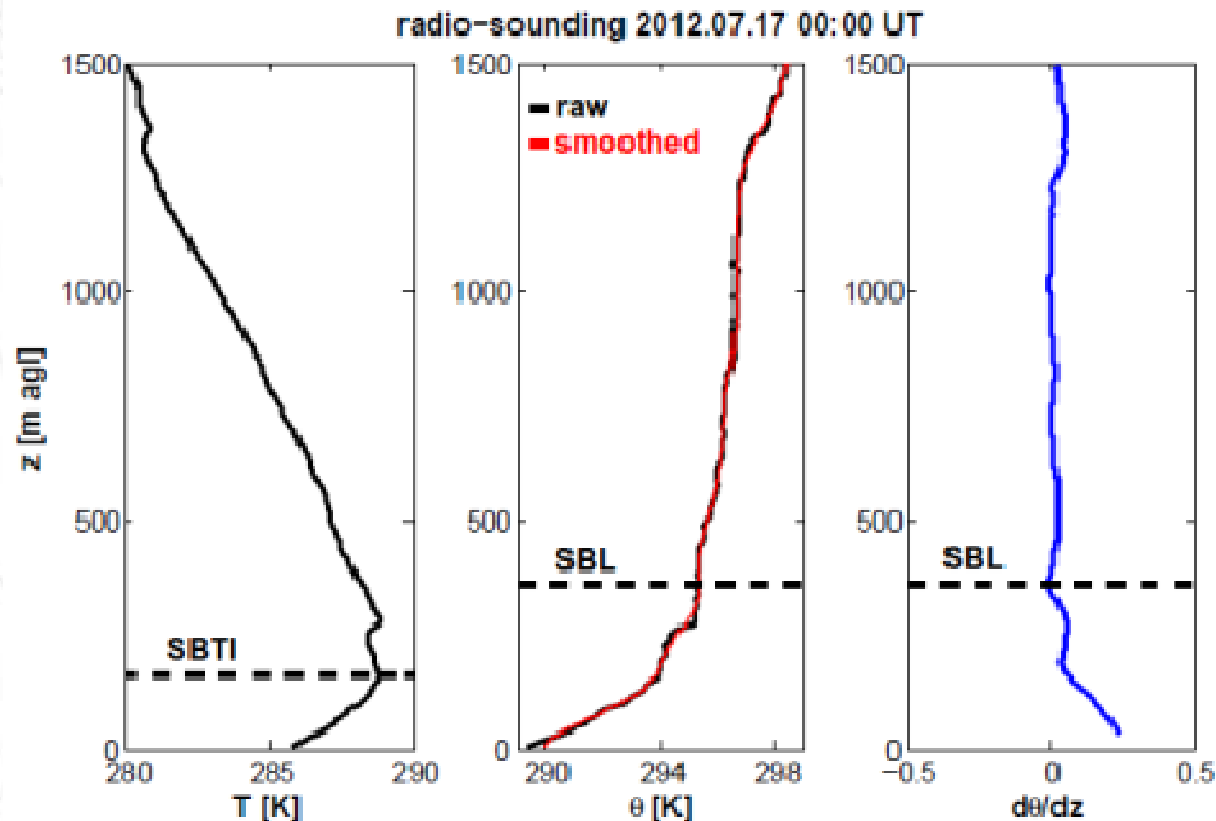
\* Holzworth, 1964; Fisher et al., 1998

# Parcel Method (PM)

- Advantages:
  - It only needs  $T$  (or  $T$  and  $H$ )  $\rightarrow \theta$  or  $\theta_v$
  - An estimate of the MLH uncertainty is obtained by varying  $T_s \pm 0.5$  K, resulting in 50-150 m
- Limitations:
  - $T_s$  has a large impact on MLH estimate (it needs a precise measurement)
  - It can only be applied to unstable conditions (Convective BL)

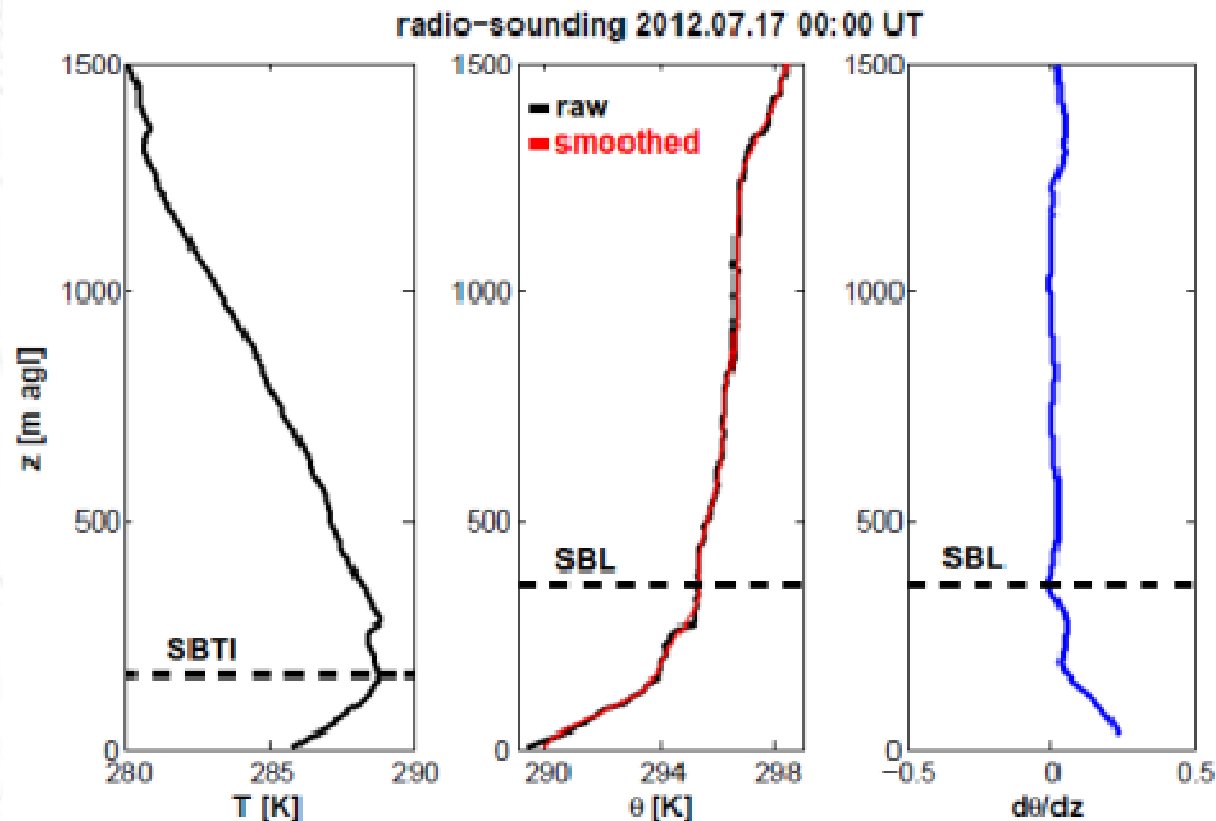
# Stable BL

- A surface-based  $T$  inversion is a clear indicator of a stable BL
- The Surface-Based Temperature Inversion (SBTI) method computes the height of the surface-based  $T$  inversion
  - where  $T$  first decreases with elevation ( $dT/dz=0$ )



# Stable BL

- Another approach is to compute the stable BL (SBL) as the height at which the gradient of  $\theta$  vanishes, i.e.  $d\theta/dz=0$
- SBL is higher than SBTI since the gradient is still positive at the height of the surface-based  $T$  inversion



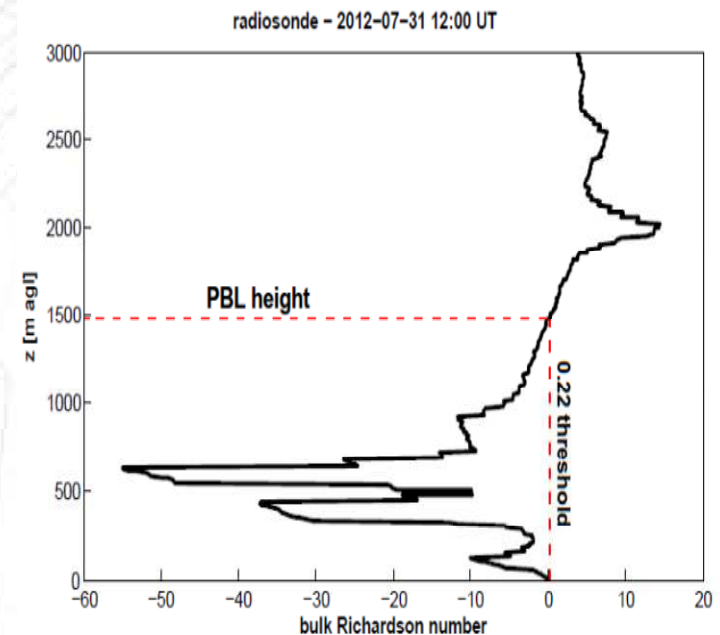


# Bulk Richardson Method

- Ri is a dimensionless number relating vertical stability and vertical wind shear (ratio of convective and wind shear produced turbul.)

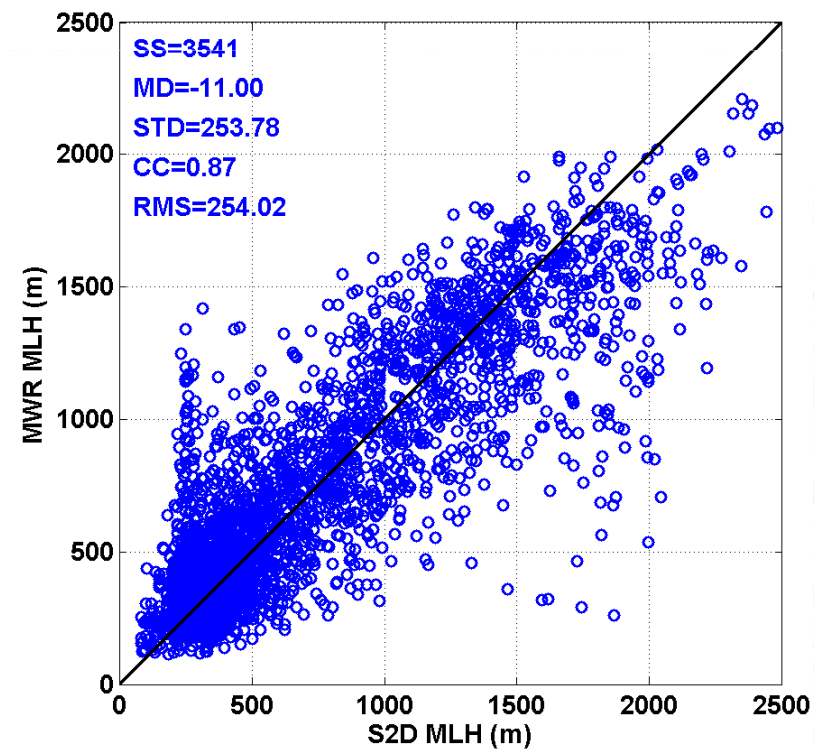
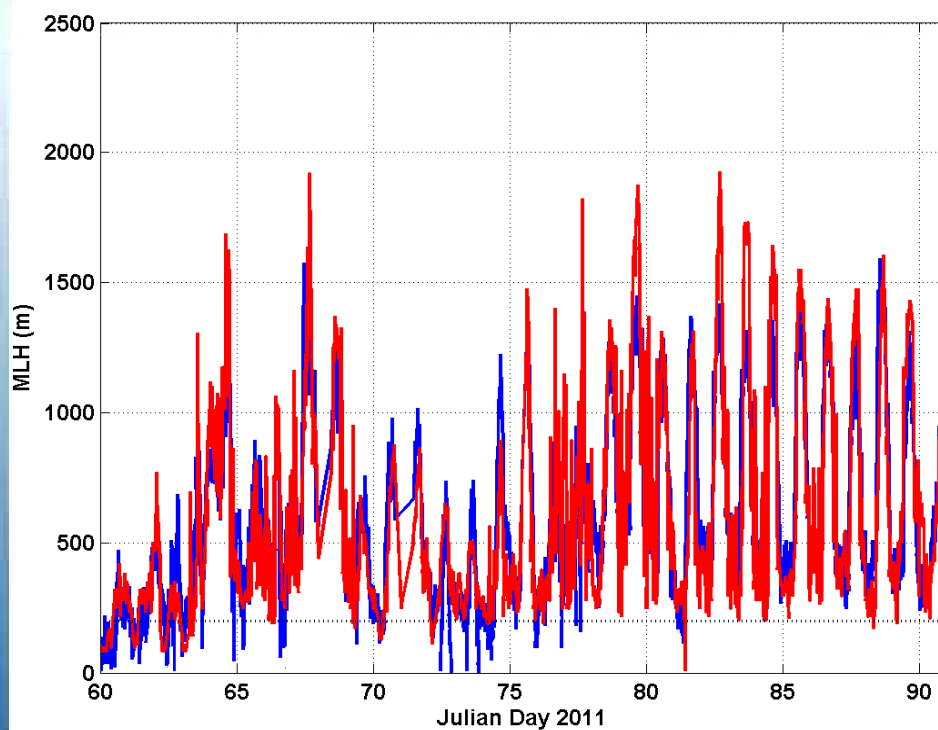
$$Ri = \frac{g\Delta\theta_v / \theta_v}{[(\Delta U)^2 + (\Delta V)^2] / \Delta Z}$$

- High values indicate unstable and/or weakly sheared conditions;
- Low values indicate weak instability and/or strong vertical shear.
- PBL height corresponds to first elevation for which Rib is greater than a critical threshold (0.22 or 0.33 for day and night conditions)



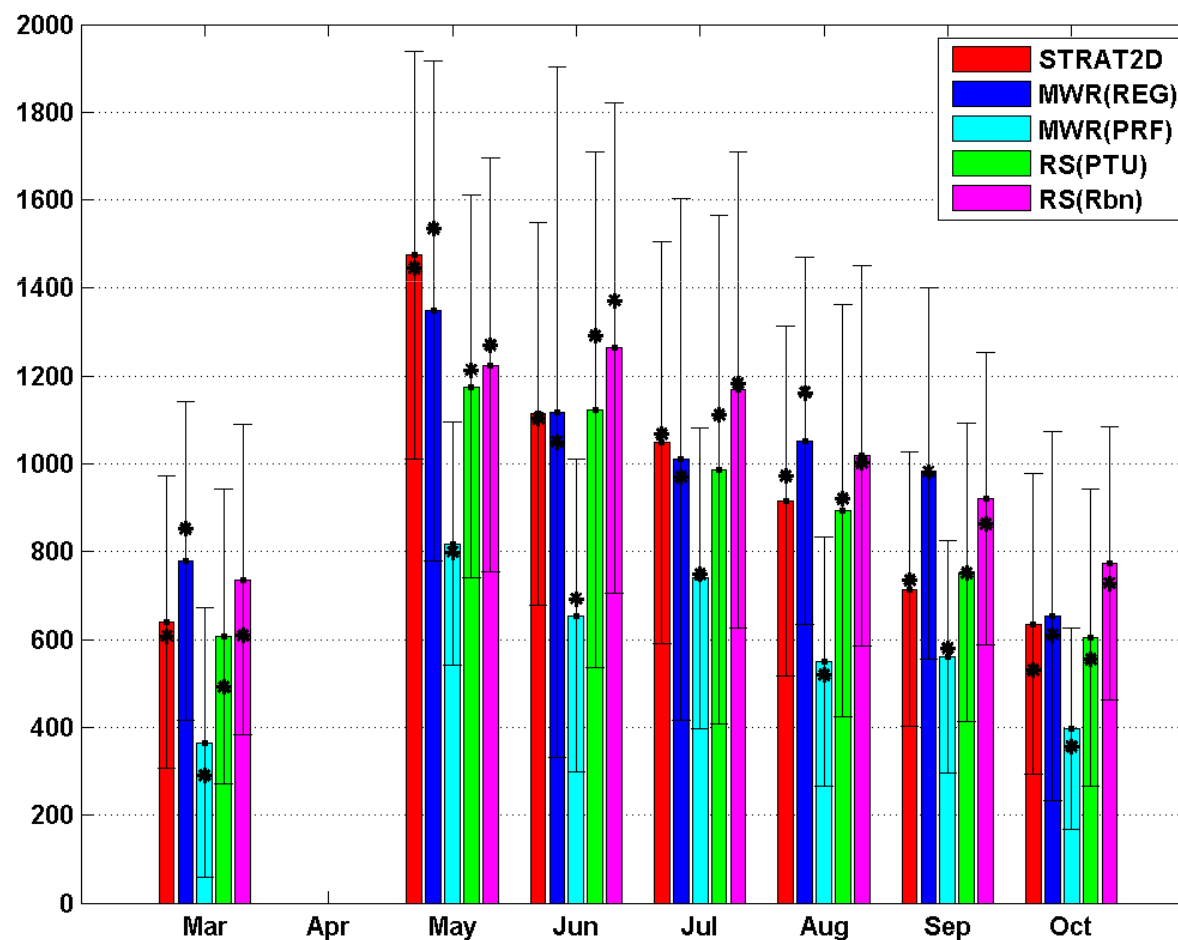
# Direct Method

- Direct estimate of MLH and SBL from MWR Tb
  - Regression or Neural Network
  - Reference for training: RS or STRA2D+ (lidar)
- Validated against reference (independent data)



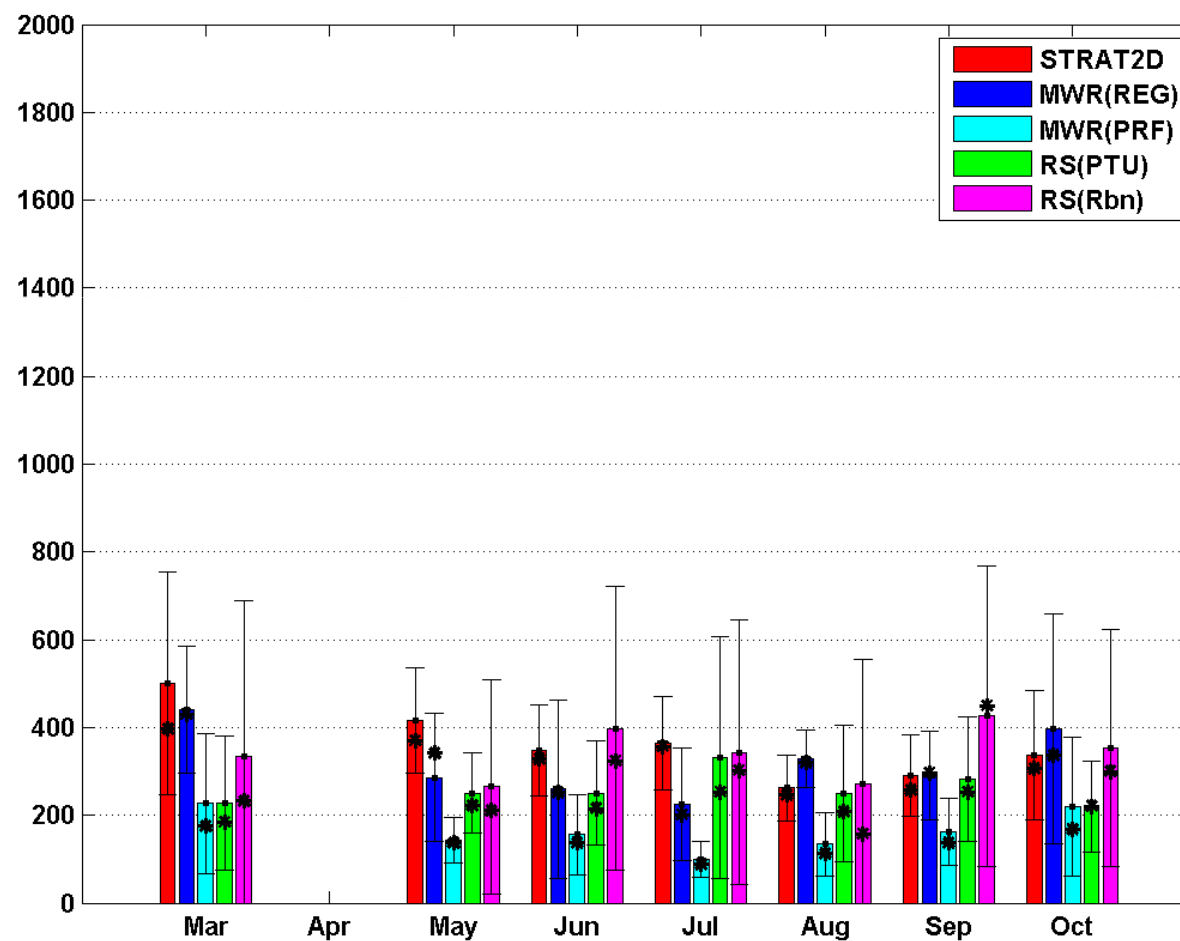
# Direct Method

## □ Convective BL



# Direct Method

## □ Stable BL



# Summary

## Summary

- Convective BL
  - Good correlations are found with other methods
  - Differences of 100–300m between various instruments and/or methods.
- Stable BL
  - MWR can provide an estimate of SBL, while other methods retrieve the residual layer (RL) height

## Bottom line

- High potential for synergy with other instruments

**Thank you very much for your attention!!**

# Relevant publications

- Cimini, De Angelis, Dupont, Pal, and Haeffelin: Mixing layer height retrievals by multichannel MWR observations, Atmos. Meas. Tech., 2013.
- Collaud Coen, M., Praz, C., Haeefe, A., Ruffieux, D., Kaufmann, P., and Calpini, B.: Determination and climatology of the planetary boundary layer height by in-situ and remote sensing methods as well as the COSMO model above the Swiss plateau, Atmos. Chem. Phys. Discuss., 14, 15419-15462, doi:10.5194/acpd-14-15419-2014, 2014.
- Seibert, P., Beyrich, F., Gryning, S. E., Joffre, S., Rasmussen, A., and Tercier, P.: Review and intercomparison of operational methods for the determination of the mixing height, Atmos. Environ., 34, 1001–1027, 2000.
- Seidel, D. J., Ao, C. O., and Li, K.: Estimating climatological planetary boundary layer heights from radiosonde observations: Comparison of methods and uncertainty analysis, J. Geophys. Res., 115, D16113, doi:10.1029/2009JD013680, 2010.

# Potential Temperature

The **potential temperature** of a air parcel at pressure  $P$  is the temperature that the parcel would acquire if adiabatically brought to a standard reference pressure  $P_0$ , usually 1000 mb.

$$\theta = T \left( \frac{P_0}{P} \right)^{R/c_p},$$

where  $T$  is the absolute temperature (in K) of the parcel,  $R$  is the gas constant of air, and  $c_p$  is the specific heat capacity at a constant pressure.